

Multi-dimensional representation of the electron density from satellite data

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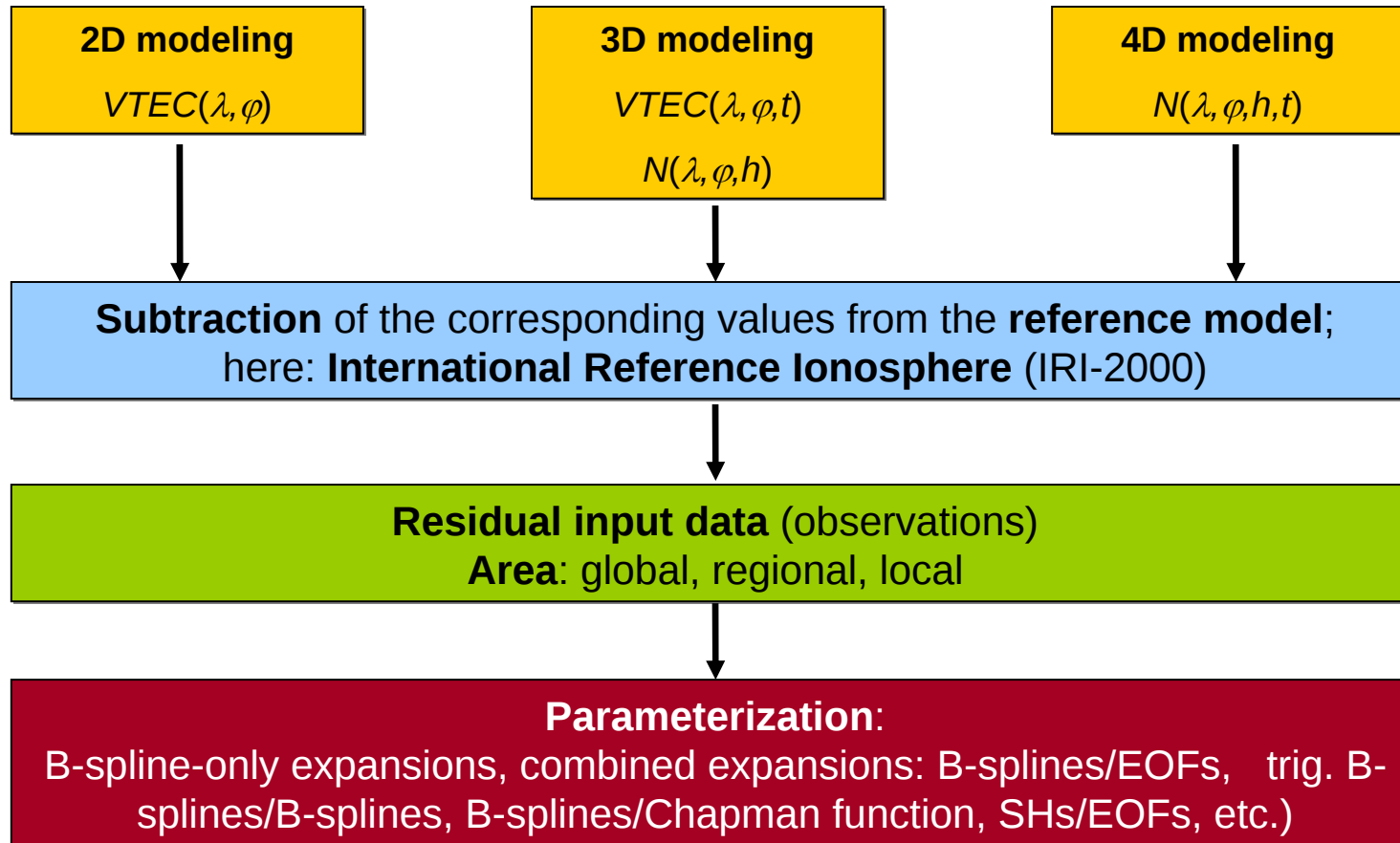


Introduction

- In the last decade **satellite missions** have become a promising tool for **measuring** ionospheric parameters such as the **electron density** or the **vertical total electron content** (VTEC) .
- **Dual-frequency GPS observations** can be used to determine the **slant total electron content** (STEC), i.e. the integral of the electron density along the ray-path of the signal.
- The insensitivity of ground-based observations to the **radial geometry** can be overcome by considering measurements from **space-borne receivers** flying on low-Earth-orbiting (LEO) satellites, e.g., the **FORMOSAT-3/COSMIC** 6-satellite constellation, **CHAMP** or **GRACE**.
- In the following, we present a general **multi-dimensional approach** for modeling **spatio-temporal variations** of the ionosphere.

Procedure (multi-dimensional input options)

- Our **multi-dimensional approach** allows the following input strategies:



- Other **ionospheric parameters** like the maximum value of the electron density $N_m F_2$ can be modeled in the same manner.

GNSS - observation equation

- The observation equation of the **geometry-free linear combination** reads

$$\begin{aligned}\phi_4(t) &= \phi_1(t) - \phi_2(t) \\ &= \alpha \cdot STEC(t) + \beta_R + \beta_S + \beta_{S,R} + e(t)\end{aligned}$$

$\beta_R, \beta_S, \beta_{S,R}$ = inter-frequency delays in receiver R and satellite S , ambiguity term ,

$\alpha, e(t)$ = constant, measurement error ,

$\phi_1(t), \phi_2(t)$ = phase observations on frequencies $f_1 = 1.5$ GHz and $f_2 = 1.2$ GHz ,

$STEC(t)$ = **Slant Total Electron Content** .

- The ambiguity term $\beta_{S,R}$ can be determined in a **pre-processing step**.
- The inter-frequency delays β_R and β_S can be determined in a **common adjustment** together with the parameters of the electron density approach we will present in the following.

STEC modeling

- The STEC is defined as the integral of the space- and time-dependent four-dimensional **electron density** $N(\mathbf{x}, t)$ along the ray-path between the satellite S and the receiver R , i.e.

$$STEC(t) = \int_R^S N(\mathbf{x}, t) ds , \quad (1)$$

wherein

$$\mathbf{x} = r \left[\cos \varphi \cos \lambda, \cos \varphi \sin \lambda, \sin \varphi \right]^T = \text{geocentric position vector}$$

λ, φ, r = longitude, latitude, radial distance , t = time

$$r = \rho + h , \quad \rho, h = \text{mean Earth radius, height} .$$

- We decompose the electron density into a **given reference model** $N_{ref}(\mathbf{x}, t)$, computed from IRI, and an **unknown correction term** $\Delta N(\mathbf{x}, t)$, i.e.

$$N(\mathbf{x}, t) = N_{ref}(\mathbf{x}, t) + \Delta N(\mathbf{x}, t) . \quad (2)$$

STEC modeling — 2

- Inserting Eq. (2) into Eq. (1) yields

$$STEC(t) = STEC_{ref}(t) + \int_R^S \Delta N(\mathbf{x}, t) ds.$$

- Approach: **3D B-splines**, i.e., a series expansion

$$\Delta N(\mathbf{x}, t) = \sum_{k_1=0}^{K_1-1} \sum_{k_2=0}^{K_2-1} \sum_{k_3=0}^{K_3-1} d_{k_1, k_2, k_3}(t) \phi_{k_1}^{J_1}(\lambda) \phi_{k_2}^{J_2}(\varphi) \phi_{k_3}^{J_3}(h)$$

in 3D products of 1D **scaling functions** $\phi_k^J(\cdot)$ with J = level (scale) and k = shift.

GNSS - observation equation — 2

- The substitution of the series expansion for $\Delta N(x, t)$ into the "STEC" equation yields

$$STEC(t) = STEC_{ref}(t) + \sum_{k_1=0}^{K_1-1} \sum_{k_2=0}^{K_2-1} \sum_{k_3=0}^{K_3-1} d_{k_1, k_2, k_3}(t) k_{k_1, k_2, k_3}(R, S),$$

wherein

$$k_{k_1, k_2, k_3}(R, S) = \int_R^S \phi_{k_1}^{J_1}(\lambda) \phi_{k_2}^{J_2}(\varphi) \phi_{k_3}^{J_3}(h) ds.$$

- The final observation equation for the geometry-free GPS observation $\phi_4(t)$ reads

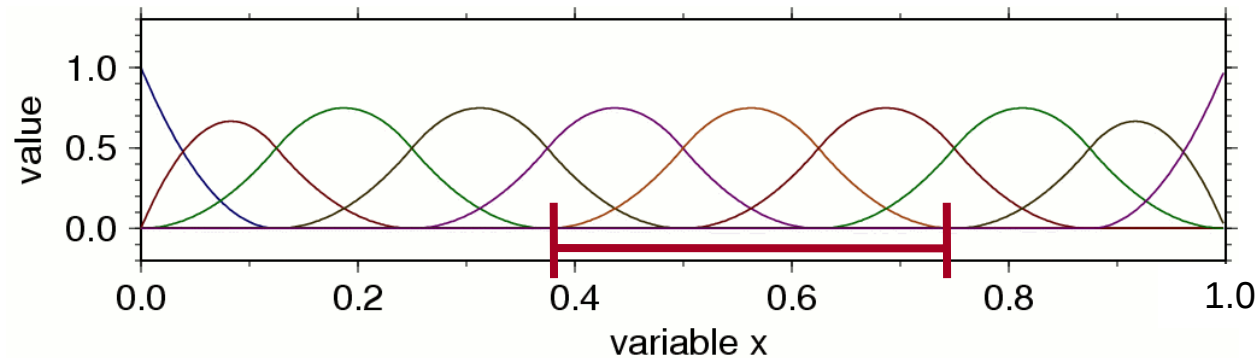
$$\begin{aligned} \phi_4(t) - \beta_{S,R} - \alpha \cdot STEC_{ref}(t) \\ = \alpha \cdot \sum_{k_1=0}^{K_1-1} \sum_{k_2=0}^{K_2-1} \sum_{k_3=0}^{K_3-1} d_{k_1, k_2, k_3}(t) k_{k_1, k_2, k_3}(R, S) + \beta_R + \beta_S + e(t). \end{aligned}$$

unknown scaling coefficients

unknown delays

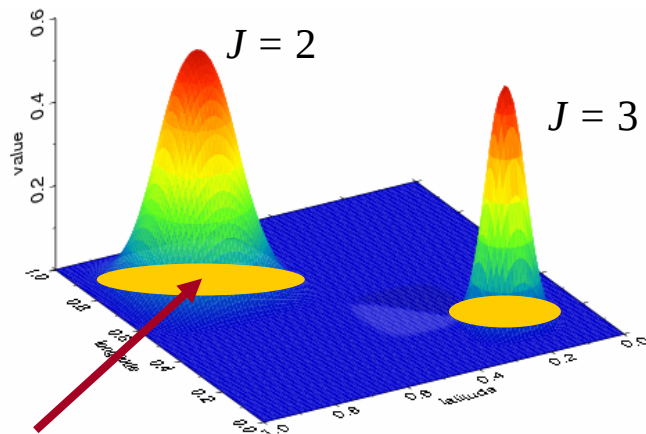
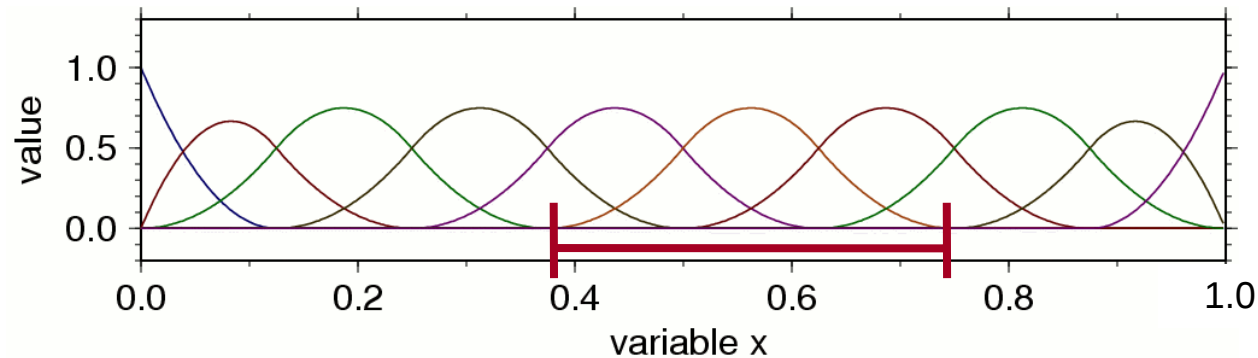
B-spline modeling

- As 1D scaling functions $\phi_k^J(x)$ of level J we choose **normalized endpoint-interpolating quadratic B-spline functions** shown in the figure below for level $J = 3$. ($\# = 2^J + 2$)



B-spline modeling

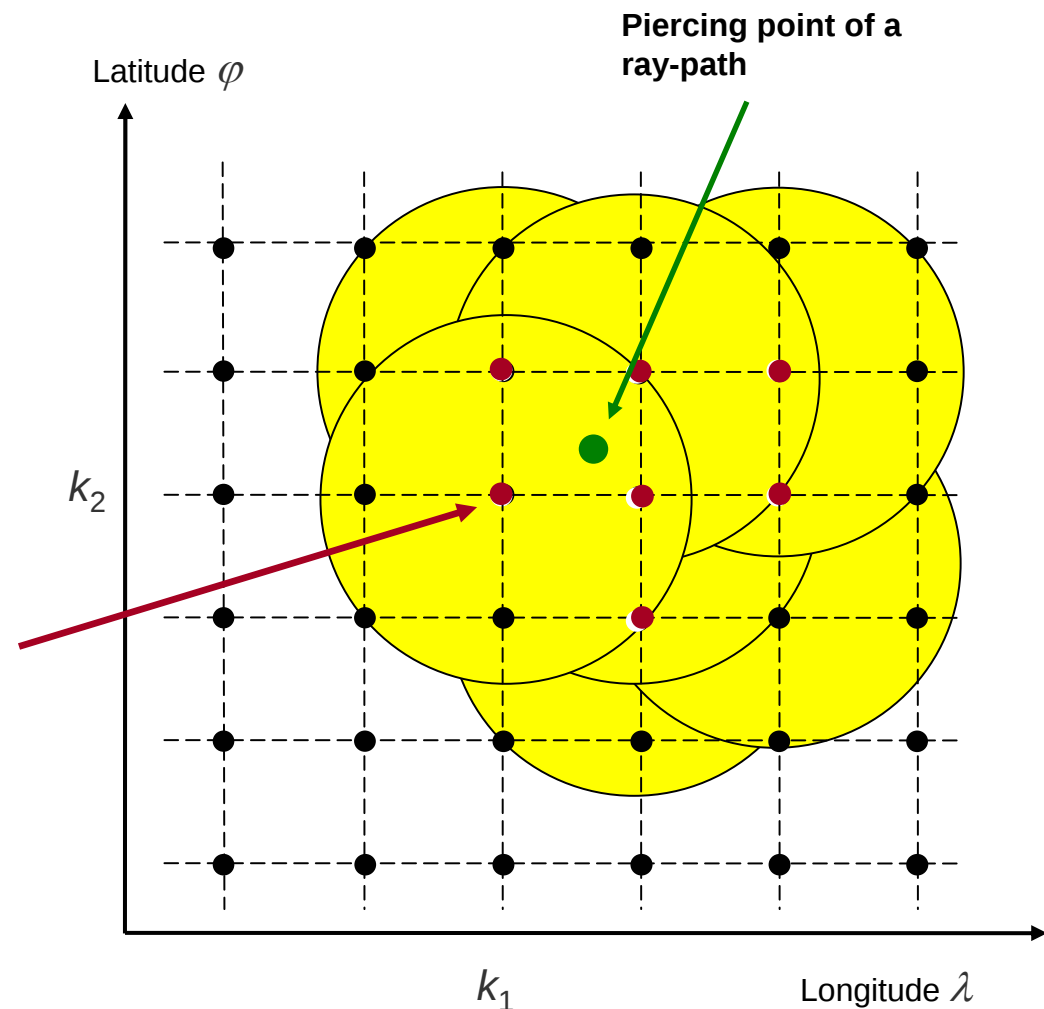
- As 1D scaling functions $\phi_k^J(x)$ of level J we choose **normalized endpoint-interpolating quadratic B-spline functions** shown in the figure below for level $J = 3$. ($\# = 2^J + 2$)



- The figure on the left-hand side demonstrates the **localizing feature** of the 2D B-splines $\phi_{k_1}^{J_1}(x) \phi_{k_2}^{J_2}(y)$, since the higher the level value J is chosen the sharper is the peak.
- Mathematically spoken B-splines are **compactly supported**. The „**non-zero**“ **zone** (influence zone, **foot print**) in the (x,y)-plane is similar to a **circle** or an **ellipse**.
- In the 3D case these non-zero zones are similar to **spheres** or **ellipsoids** (biaxial or triaxial).

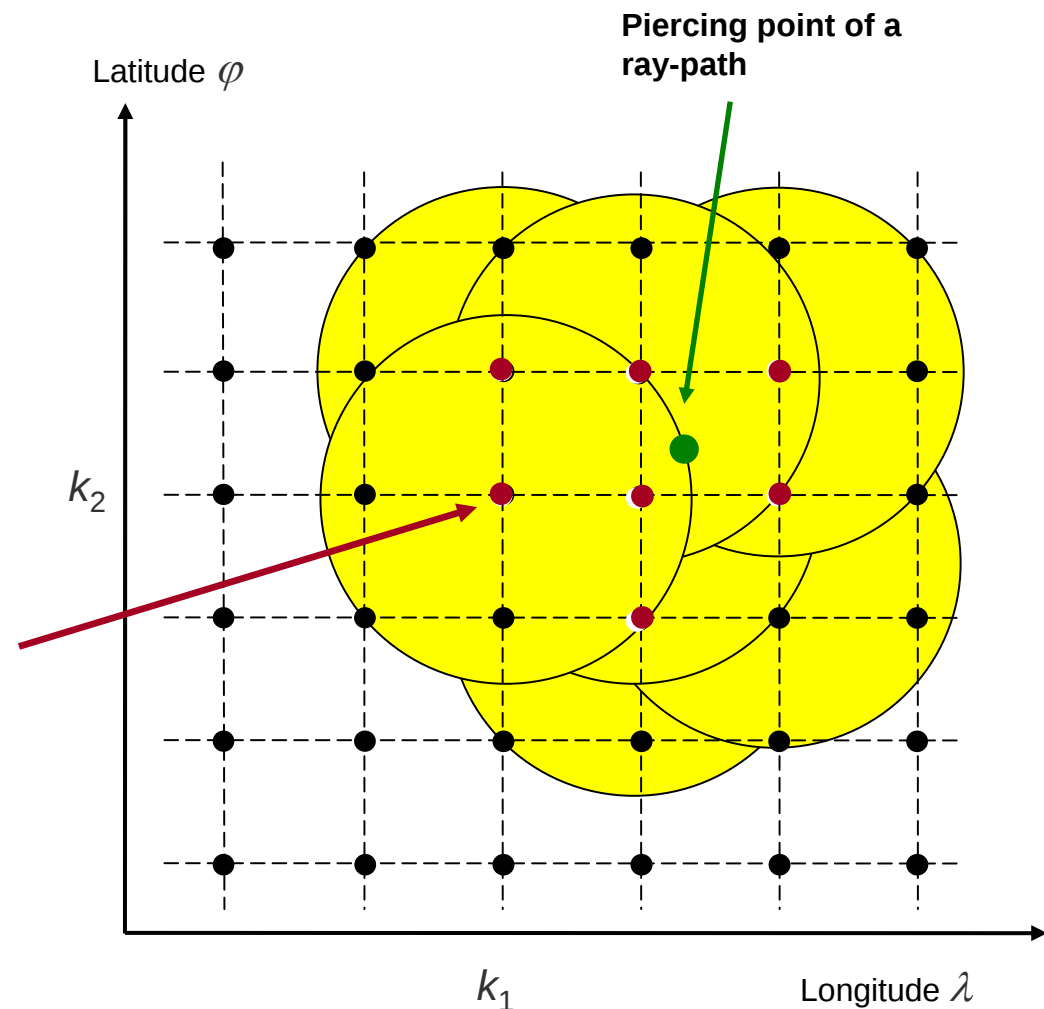
B-spline modeling — 2

- The centers (k_1, k_2) of **the foot prints** of the 2D B-spline functions $\phi_{k_1}^{J_1}(\lambda) \phi_{k_2}^{J_2}(\varphi)$ define a **regular grid** within the unit square.
- Due to their compact support (yellow **foot prints**) only B-splines related to the red dots, have **non-zero entries** in the observation equation shown before.
- The **series (scaling) coefficient** d_{k_1, k_2} can be determined by the GNSS observation related to the piercing point (green dot).

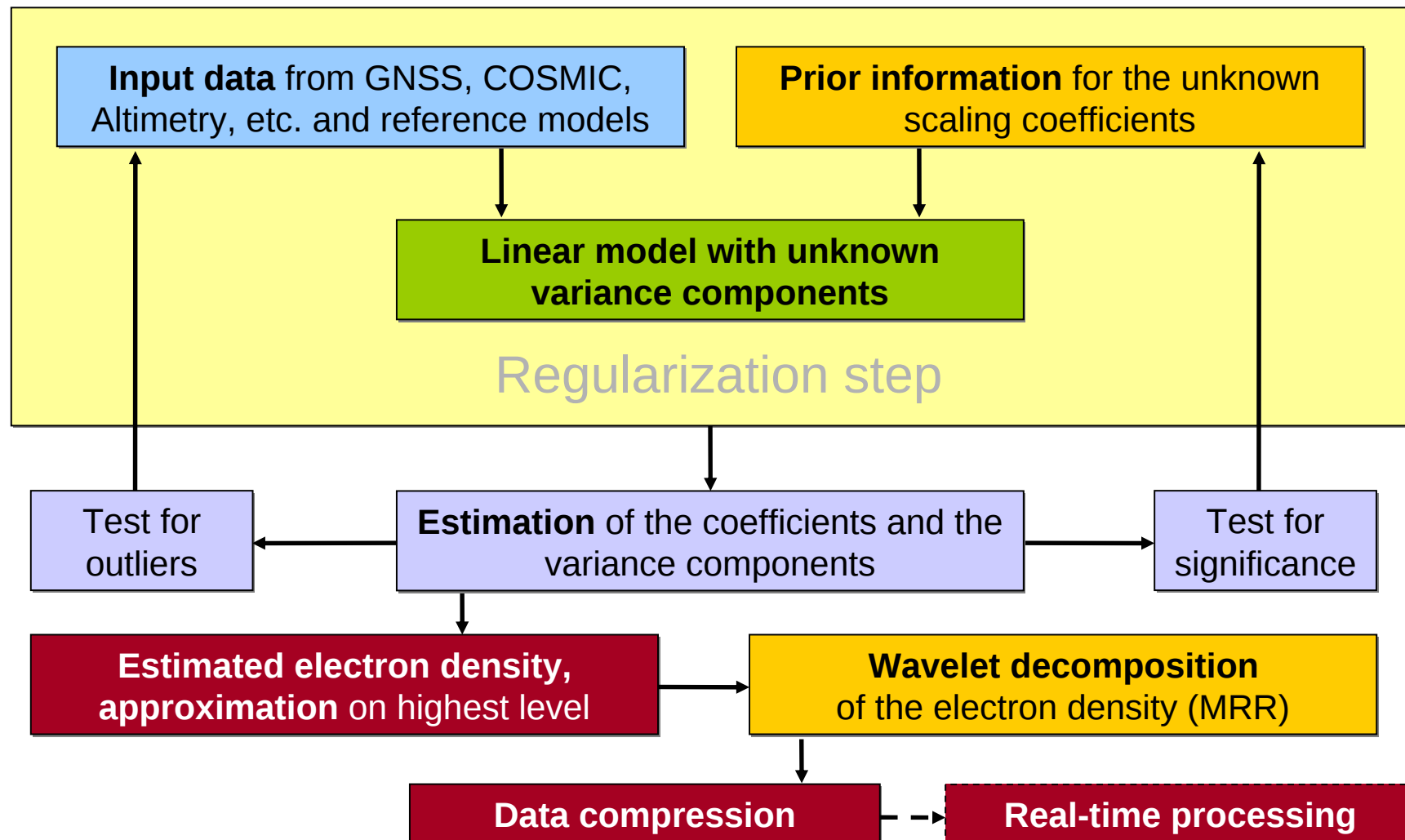


B-spline modeling — 2

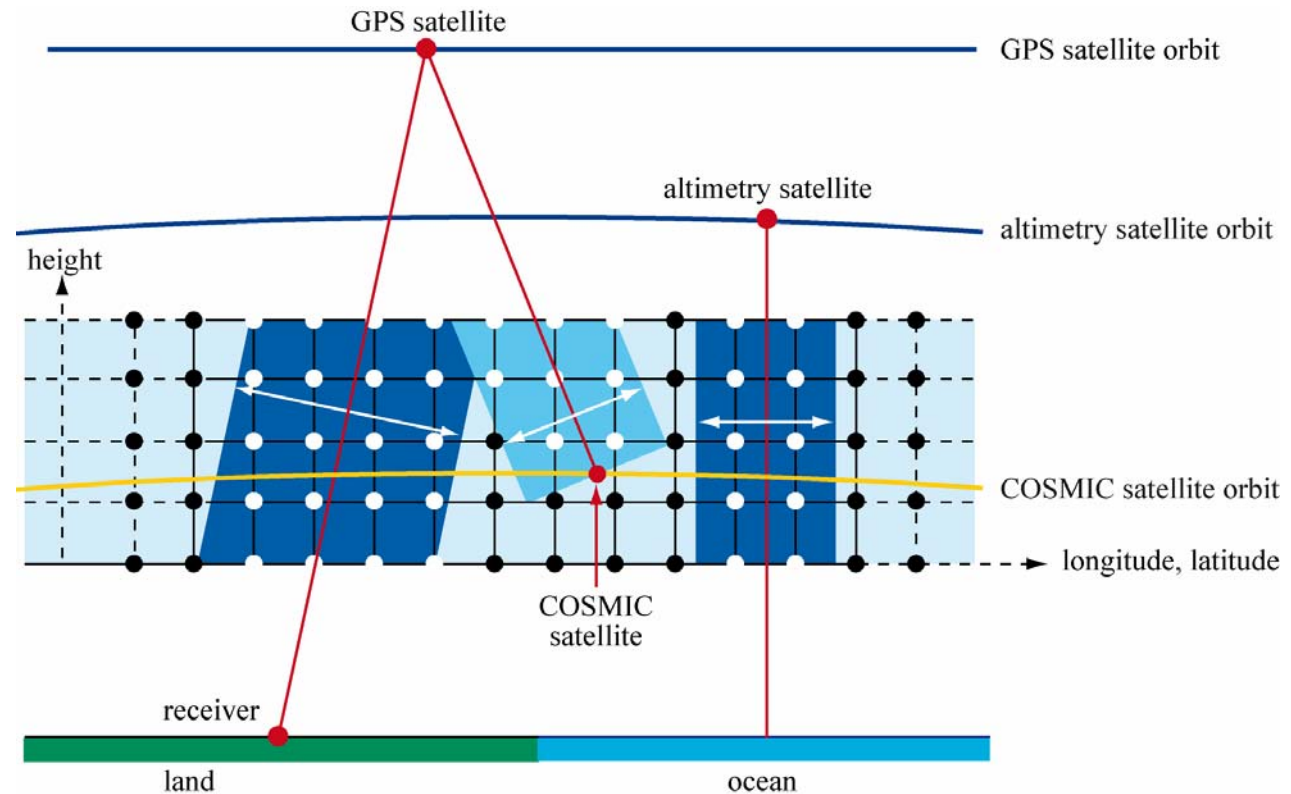
- The centers (k_1, k_2) of **the foot prints** of the 2D B-spline functions $\phi_{k_1}^{J_1}(\lambda) \phi_{k_2}^{J_2}(\varphi)$ define a **regular grid** within the unit square.
- Due to their compact support (yellow **foot prints**) only B-splines related to the red dots, have **non-zero entries** in the observation equation shown before.
- The **series (scaling) coefficient** d_{k_1, k_2} can be determined by the GNSS observation related to the piercing point (green dot).
- If the location of the observation lies close to the border of the foot print the coefficient is weakly determined → **prior information** for d_{k_1, k_2} .



Procedure (adjustment, MRR, data compression)



Input data



- The relative **GPS-LEO geometry** is changing rapidly (including **occultation** measurements).
- LEO measurements **stabilize** the estimation process.
- Dual-frequency **altimetry** satellites measure always the **vertical electron content** (more or less only over the oceans).

Combination of input data

- For each of the **different observation types** (including altimetry and terrestrial measurements) it is possible to derive an **observation equation** for estimating the unknowns of the electron density B-spline model within a specific spatio-temporal (4D) region $\mathbb{B}^4$.
- We assume now that for each of the altogether p techniques an **observation vector** $\mathbf{y}_i$, $i = 1, \dots, p$ is available, e.g. $\mathbf{y}_1$ may contain all reduced GNSS geometry-free observations measured from ground stations within $\mathbb{B}^4$.
- Hence, for each observation vector $\mathbf{y}_i$ a **Gauss-Markov model**

$$\mathbf{y}_i + \mathbf{e}_i = \mathbf{A}_i \boldsymbol{\beta} \quad \text{with} \quad D(\mathbf{y}_i) = \sigma_{y,i}^2 \mathbf{P}_{y,i}^{-1}$$

can be established.

- The quantities $\sigma_{y,i}^2$ and $\mathbf{P}_{y,i}$ are the unknown variance factor and the given positive definite weight matrix.
- The vector $\boldsymbol{\beta}$ consists of the unknown **B-spline scaling coefficients** $d_{k_1,k_2,k_3}(t)$, the unknown delays β_R and β_S and other auxiliary parameters. Note, the scaling coefficients have to be discretized w.r.t. the time.

Combination of input data — 2

- With the **prior information** for the expectation vector $E(\boldsymbol{\beta}) = \boldsymbol{\mu}_\beta$ and the covariance matrix $D(\boldsymbol{\beta}) = \mathbf{P}_\beta^{-1}$ for the unknown B-spline scaling coefficients the additional linear model

$$\boldsymbol{\mu}_\beta + \mathbf{e}_\beta = \boldsymbol{\beta} \quad \text{with} \quad D(\boldsymbol{\mu}_\beta) = \sigma_\beta^2 \mathbf{P}_\beta^{-1}$$

can be formulated; $\mathbf{e}_\beta$ = error vector of the prior information, σ_β^2 = unknown variance factor.

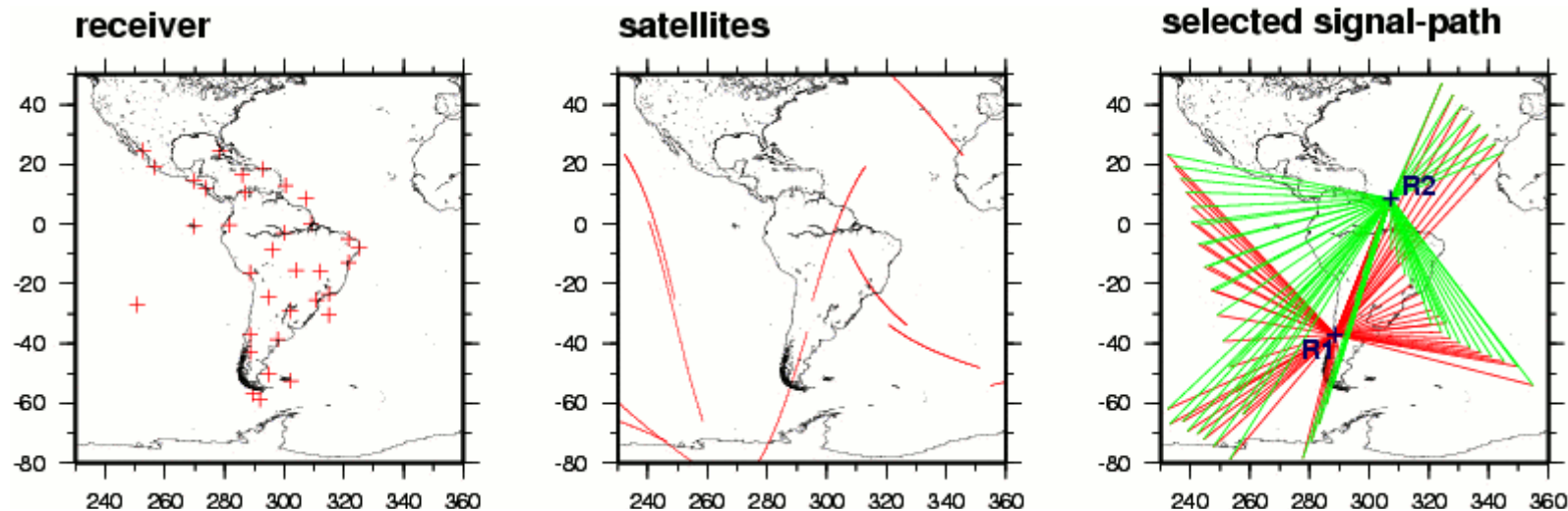
- The combination of the models yields the **normal equation system**

$$\left(\sum_{i=1}^p \frac{1}{\sigma_{y,i}^2} \mathbf{A}_i^T \mathbf{P}_{y,i} \mathbf{A}_i + \frac{1}{\sigma_\beta^2} \mathbf{P}_\beta \right) \hat{\boldsymbol{\beta}} = \sum_{i=1}^p \frac{1}{\sigma_{y,i}^2} \mathbf{A}_i^T \mathbf{P}_{y,i} \mathbf{y}_i + \frac{1}{\sigma_\beta^2} \mathbf{P}_\beta \boldsymbol{\mu}_\beta$$

for the **unknown parameter vector** $\boldsymbol{\beta}$ with **unknown variance components** $\sigma_{y,i}^2$, $i = 1, \dots, p$ and σ_β^2 .

- For solving this problem we apply a **fast Monte-Carlo implementation** of the iterative maximum-likelihood **Variance Component Estimation** (MCVCE, Koch and Kusche, 2001).

GNSS modeling



- We chose 34 stations in Central and South America as well as 69 positions along the GPS satellite orbits. Hence, we evaluate altogether $34 \cdot 69 = 2346$ simulated observations.
- The ray-paths with a zenith angle larger than 80° were neglected.
- For each measurements we have to evaluate the integrals over the 3-dimensional B-spline functions along the ray-paths. We perform this integration by **Gauss quadrature**.

For more details see:

Zeilhofer, C. et al. (2008): *Regional 4-D modeling of the ionospheric electron density from satellite data and IRI*. *Advances in Space Research (JASR)*, accepted, in press

Zeilhofer, C. (2008): *Multi-dimensional B-spline modeling of spatio-temporal ionospheric signals*. *DGK, Series A*, 123

GNSS modeling — 2

- To test our B-spline approach we simulate GNSS observations. For that purpose we consider the **residual electron density**

$$\Delta N(\mathbf{x}, t) = N_{sim}(\mathbf{x}, t) - N_{ref}(\mathbf{x}, t)$$

wherein $N_{sim}(\mathbf{x}, t)$ are simulated electron density values and $N_{ref}(\mathbf{x}, t)$ are the values of the reference model, namely a low-pass filtered IRI version.

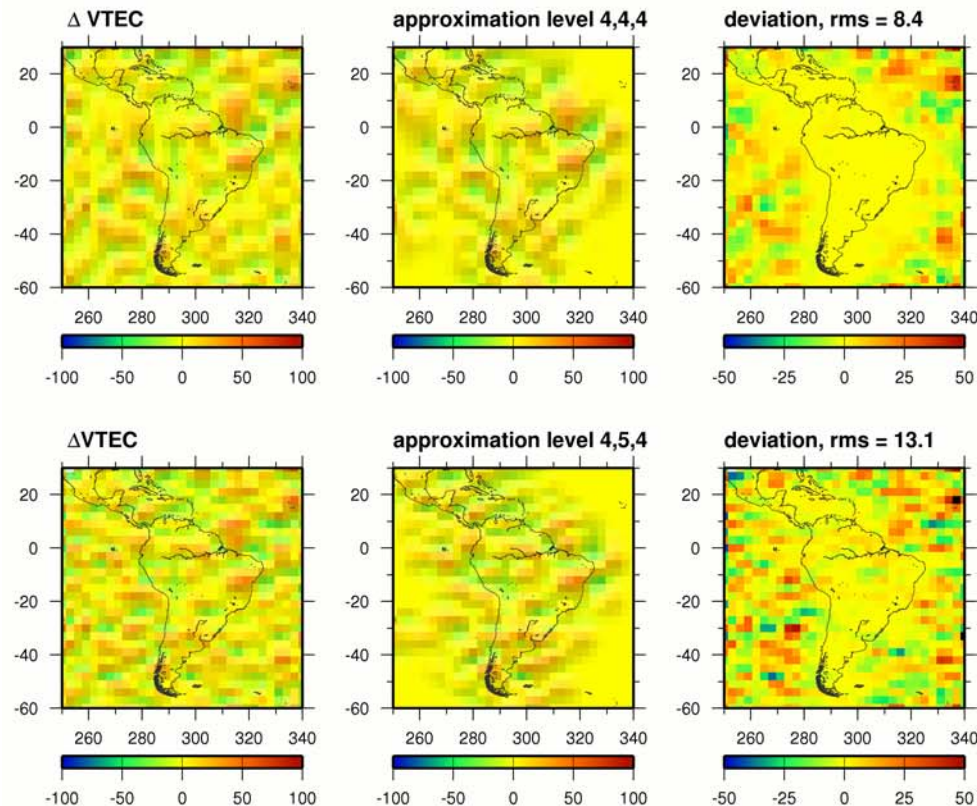
- To simulate GPS observations we set time $t = t_0$ and integrate along the ray-paths over the residual electron density $\Delta N(\mathbf{x}, t)$ and get **residual STEC** values

$$\Delta STEC(R, S) = \int_R^S \Delta N(\mathbf{x}, t) ds$$

- Replacing the integrand by the 3D series expansion for the residual electron density yields

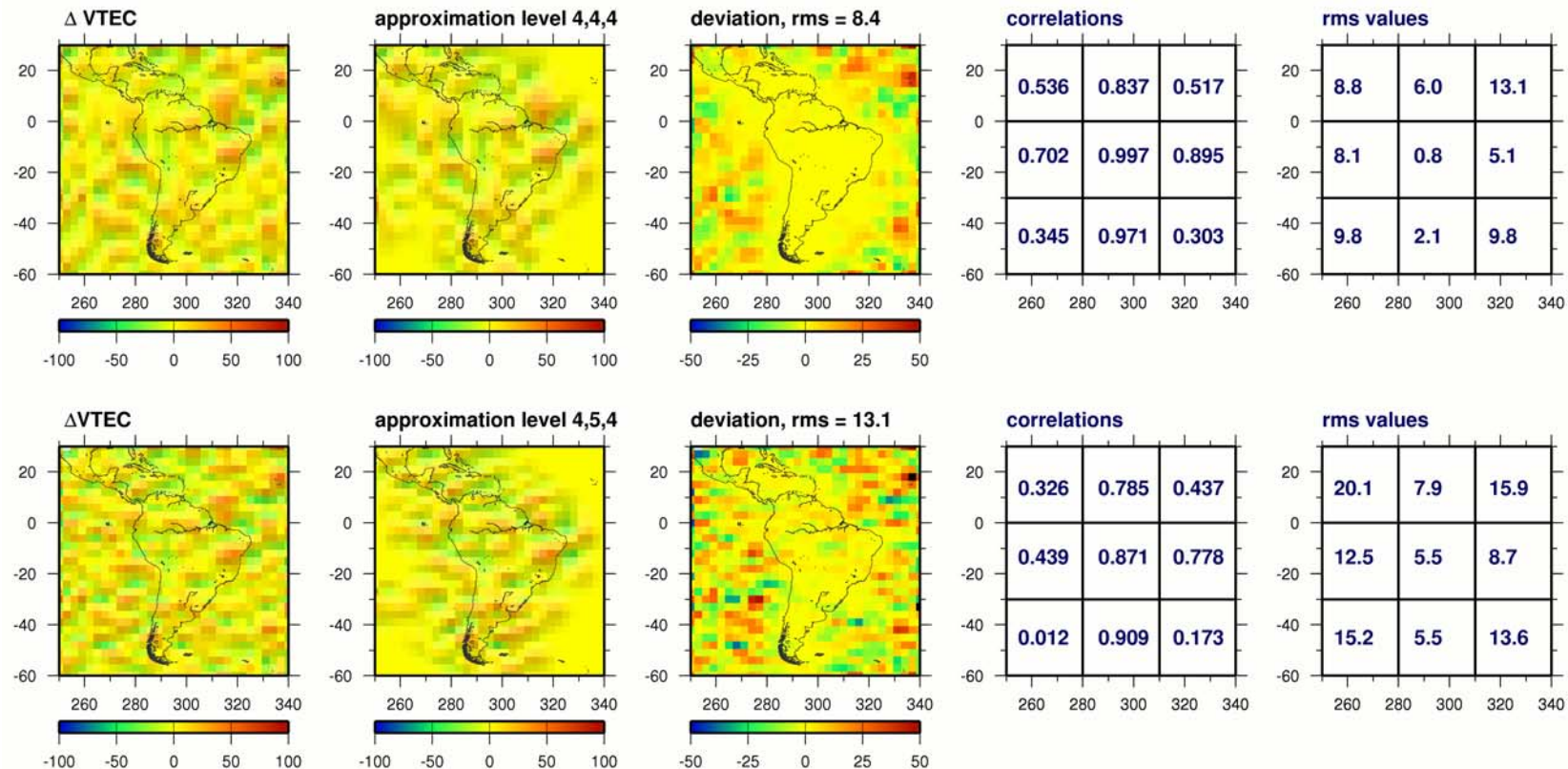
$$\Delta STEC(R, S) + e(R, S) = \sum_{k_1=0}^{K_1-1} \sum_{k_2=0}^{K_2-1} \sum_{k_3=0}^{K_3-1} d_{k_1, k_2, k_3} k_{k_1, k_2, k_3}(R, S).$$

GNSS modeling — 3



- The two rows show respectively:
 - **left panels:** 2 simulated residual VTEC data sets (bottom: higher frequency content in latitude)
 - **mid panels:** estimated B-spline models of the levels (4,4,4) and (4,5,4) respectively
 - **right panels:** deviations of the input data and the estimations
- The distribution of the data (receiver- and satellite-positions) influences the deviations.
- We divide the area into 9 subareas and calculate the correlations and rms values of the deviations.
- Over the continent we have high correlations and low rms values, this is due to the advantageous distribution of the ray-paths.

GNSS modeling — 3

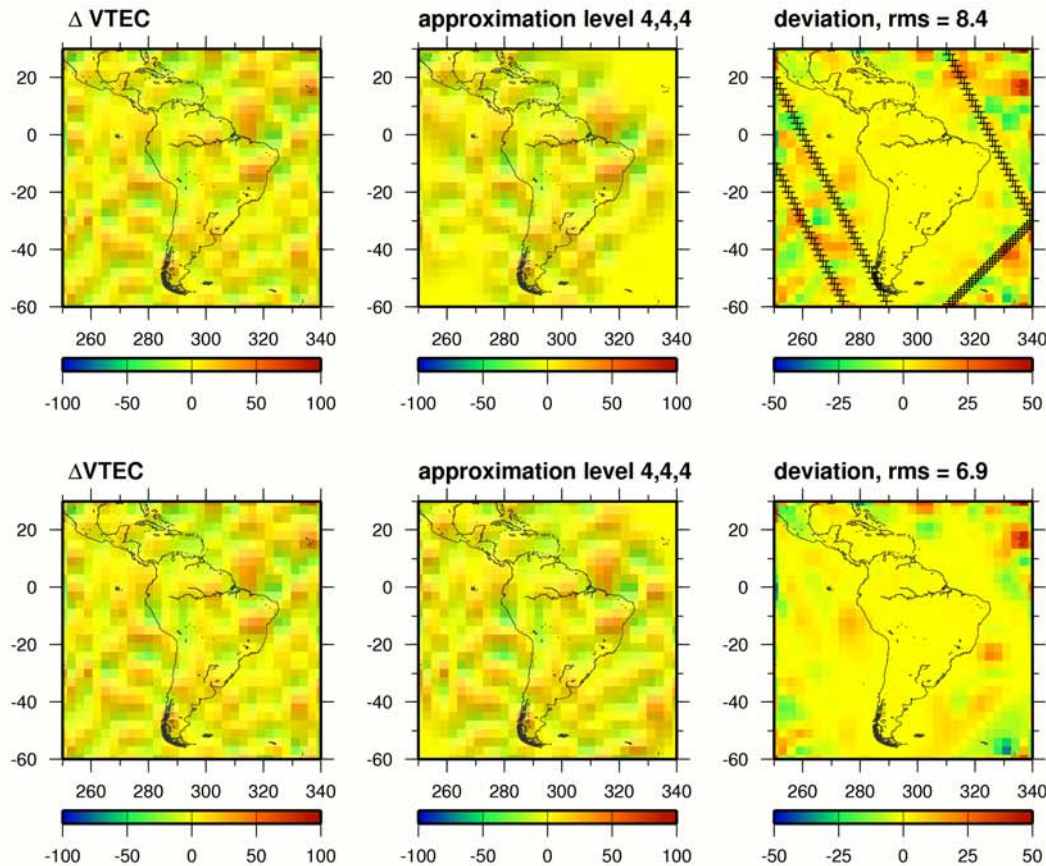


- The distribution of the data (receiver- and satellite-positions) influences the deviations.
- We divide the area into 9 subareas and calculate the correlations and rms values of the deviations.
- The combination of **altimetry** and GNSS observations combines observations that lie mainly above the ocean and above the continent.

GNSS modeling — 4

- We add an additional linear model for altimetry observations. The observation equation reads

$$\Delta VTEC(R, S) + e(R, S) = \sum_{k_1=0}^{K_1-1} \sum_{k_2=0}^{K_2-1} \sum_{k_3=0}^{K_3-1} d_{k_1, k_2, k_3} k_{k_1, k_2, k_3}(R, S).$$

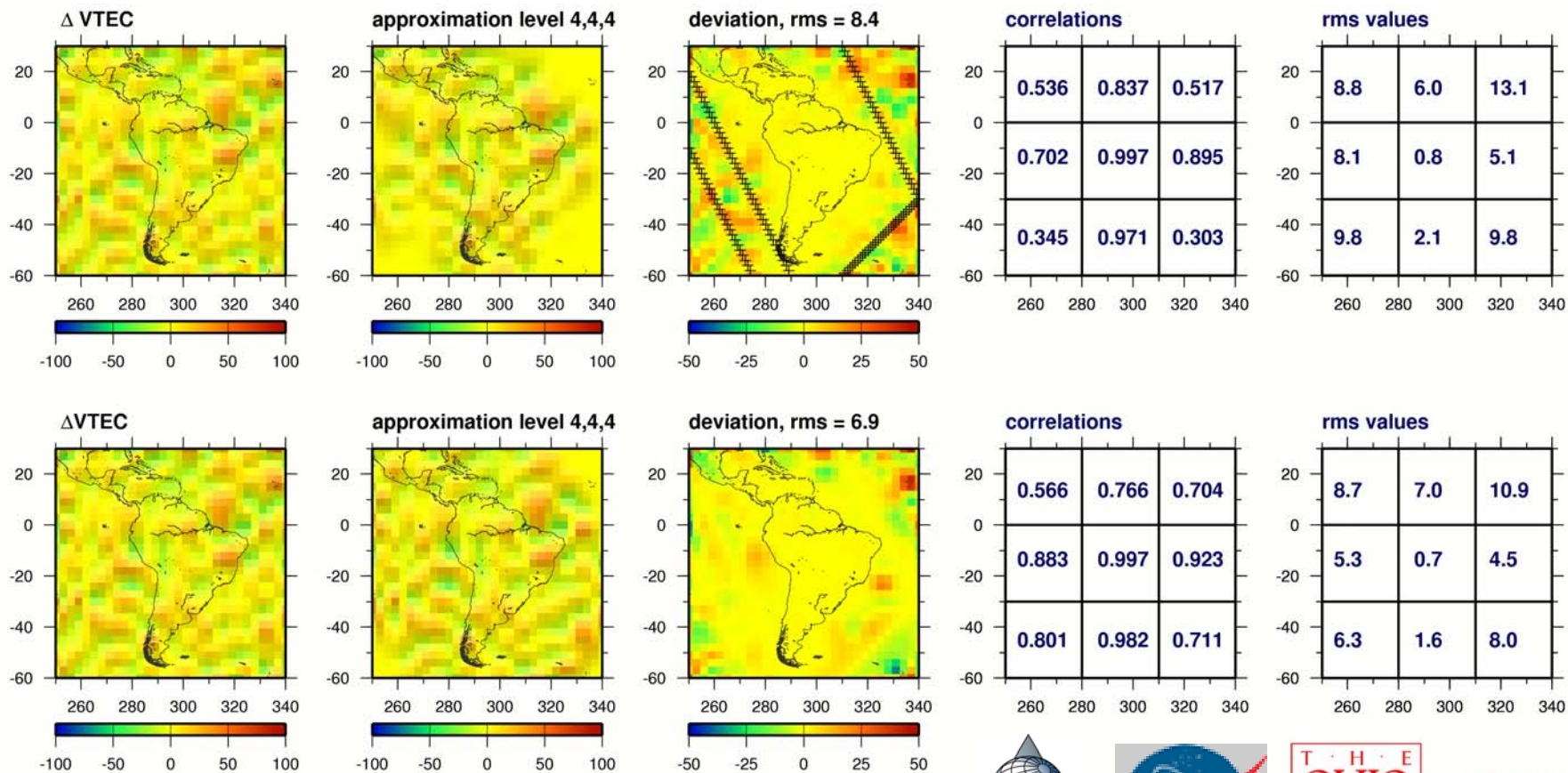


- The black crosses show the observation sites of the simulated altimetry measurements.
- The combination improves clearly the accuracy of the results over the oceans.
- The bottom panels show the combination results.
- The rms value decreases from 8.4 TECU to 6.9 TECU.

GNSS modeling — 4

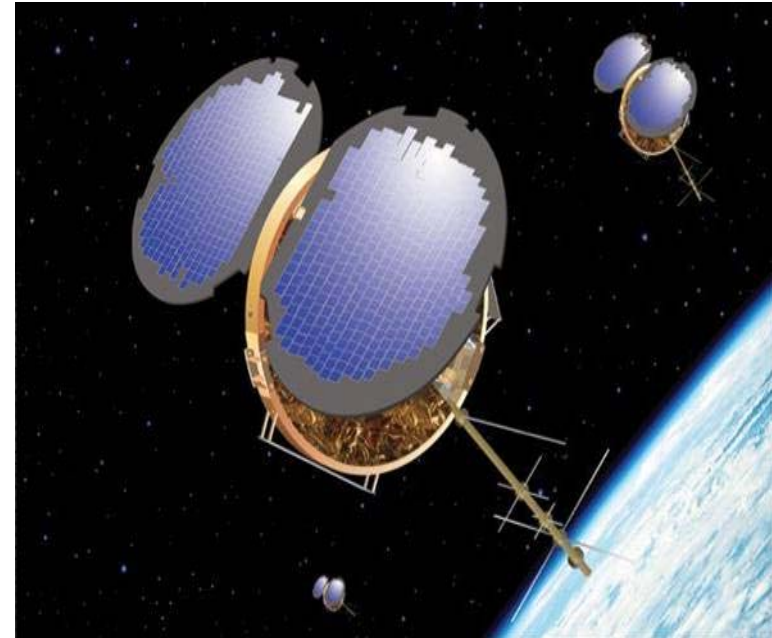
- We add an additional linear model for altimetry observations. The observation equation reads

$$\Delta VTEC(R, S) + e(R, S) = \sum_{k_1=0}^{K_1-1} \sum_{k_2=0}^{K_2-1} \sum_{k_3=0}^{K_3-1} d_{k_1, k_2, k_3} k_{k_1, k_2, k_3}(R, S).$$



COSMIC input data

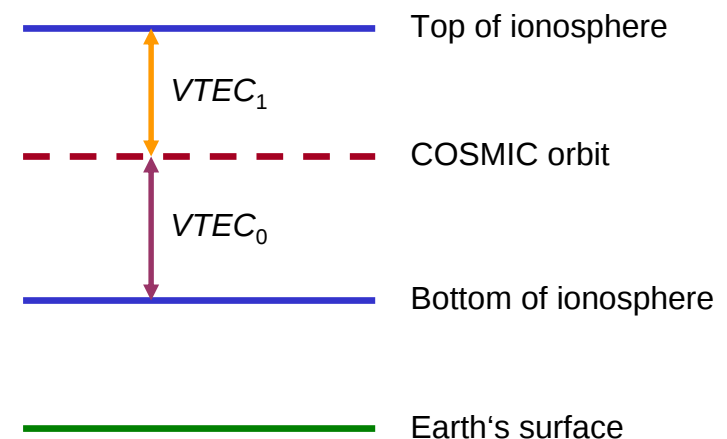
- **COSMIC/FORMOSAT-3** is the **C**onstel-
lation **O**bserving **S**ystem for **M**eteorology,
Ionosphere and **C**limate and Taiwan's
Formosa **S**atellite Mission #3, a joint
Taiwan-U.S. project.
- The six COSMIC satellites were success-
fully launched on April 14th, 2006. Over the
first year, the satellites had been gradually
boosted from their initial orbit of 400 km to
the final orbit of 800 km.
- Each spacecraft is equipped with three
instruments, namely a **GPS (occultation)**
receiver (GOX), a **tiny ionospheric pho-**
tometer (TIP) and a **tri-band beacon**
(TBB).



from: L.-C. Tsai, C.H. Liu, and T.T. Hsiao

COSMIC input data — 2

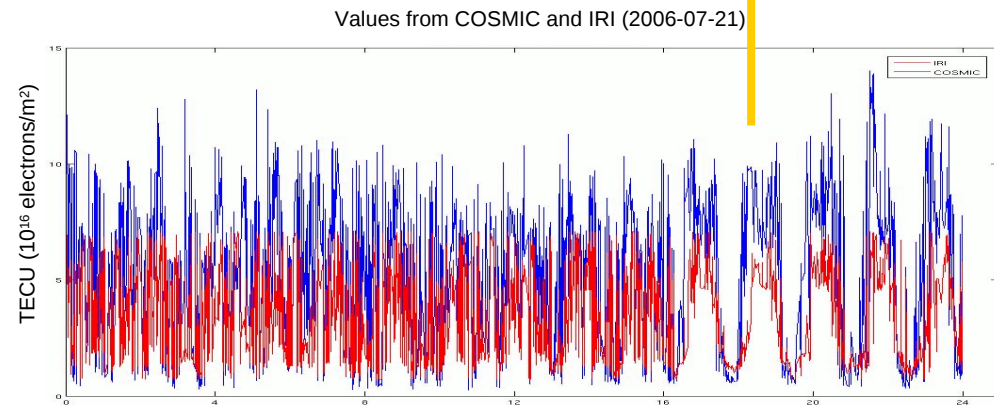
- In our investigation we calculate the electron density in a pre-processing step from so-called **compensated STEC** values by an "improved" Abel transform.
- To be more specific, we do not assume locally spherical symmetry within the ionosphere, but consider the effect of **large-scale horizontal gradients** and/or **inhomogeneous electron density distribution**.
- For more details see the paper: Tsai, Lung-Chih and Wei-Hsiung Tsai (2004): Improvement of GPS/MET ionospheric profiling and validation using the Chung-Li ionosonde measurements and the IRI model. (TAO,15(4), 589-607)
- A "**measured**" **VTEC** value (observation) mean the sum of two parts:
 - $VTEC_0$ is the integration of the calculated electron density along the vertical from bottom to orbital height.
 - $VTEC_1$ is an extrapolated value provided by Lung-Chih Tsai using a Chapman layer model.



COSMIC input data — 3

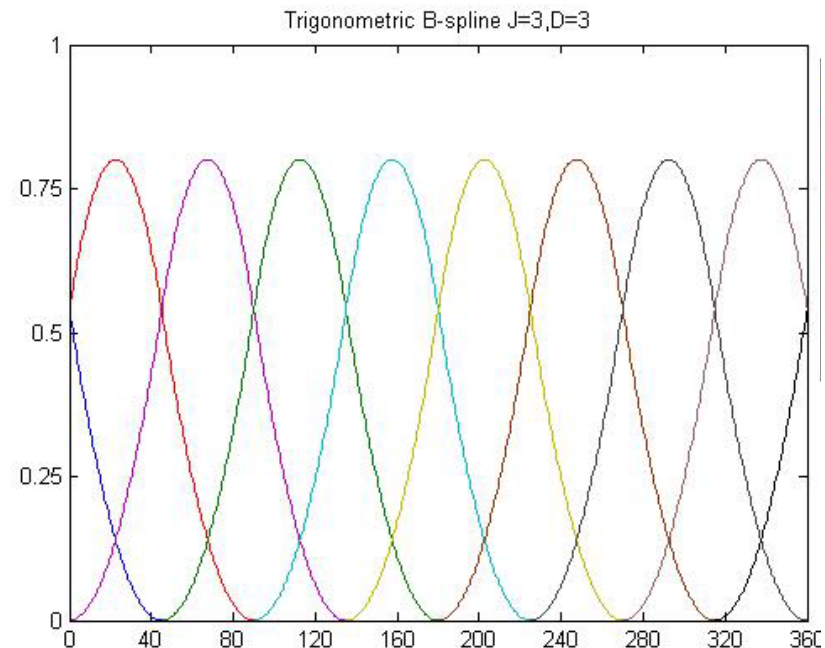
- The table shows the statistical values of the COSMIC input data and the corresponding VTEC values from IRI.
- The lower figure shows the “measured” COSMIC VTEC input data $VTEC_0$ between orbital height and the bottom of the ionosphere and the corresponding $VTEC_0$ values from IRI. (outlier detection was performed)

Edited (2547 points)		
	COSMIC	IRI
Mean	5.18	3.45
Standard Deviation	3.22	2.03
RMS	6.10	4.00
Correlation coefficient	0.76	

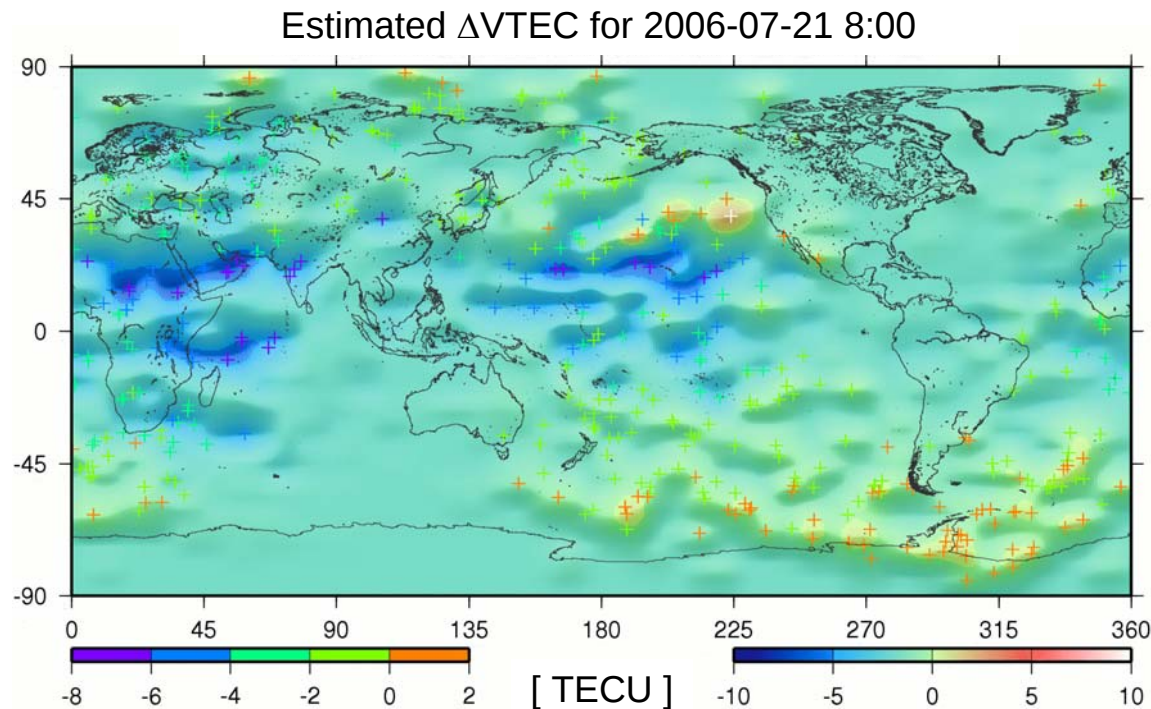


COSMIC modeling

- In case of **regional modeling** the ray-paths of **occultation** measurements might leave the region of investigation → **global model approach**.
- The figure shows the **trigonometric B-spline** functions defined on the interval between 0° to 360° .
- These functions are also **compactly supported**, i.e. the non-zero (influence) zone is finite.
- In opposite to the endpoint-interpolating B-splines the trigonometric B-splines are “**wrapping around**”.



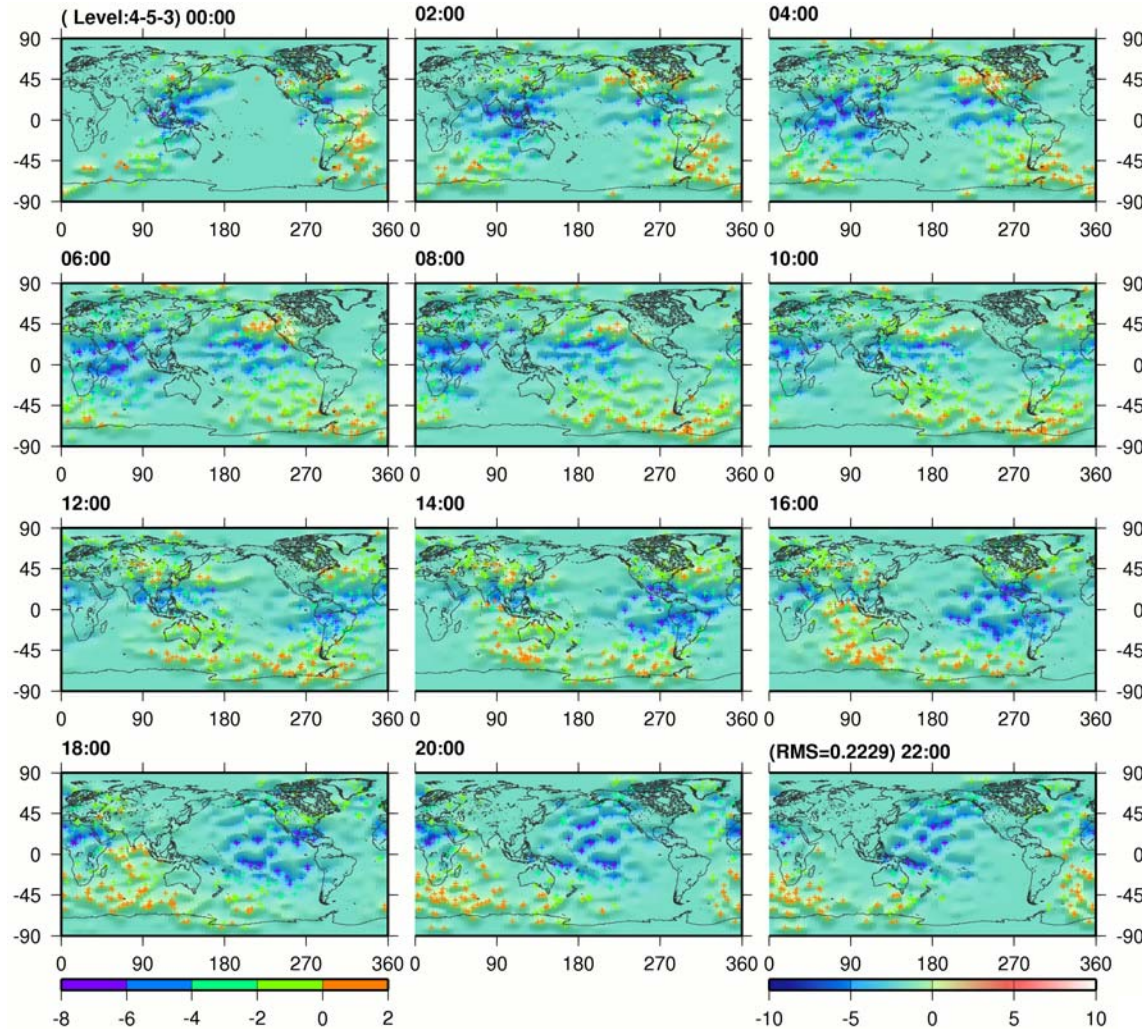
COSMIC modeling — 2



- We apply this global approach for modeling Δ VTEC, i.e. we introduce **trigonometric B-splines** in longitude (level 4).
- For modeling the latitude- and time-dependency we again use **endpoint interpolating B-splines** (level 5 and level 3, resp.).
- In the figures the crosses indicate the observation sites (**blue** and **green** colours mean negative values, **orange** positive values).
- Structures are modelable only in regions where observations are available.

COSMIC modeling — 2

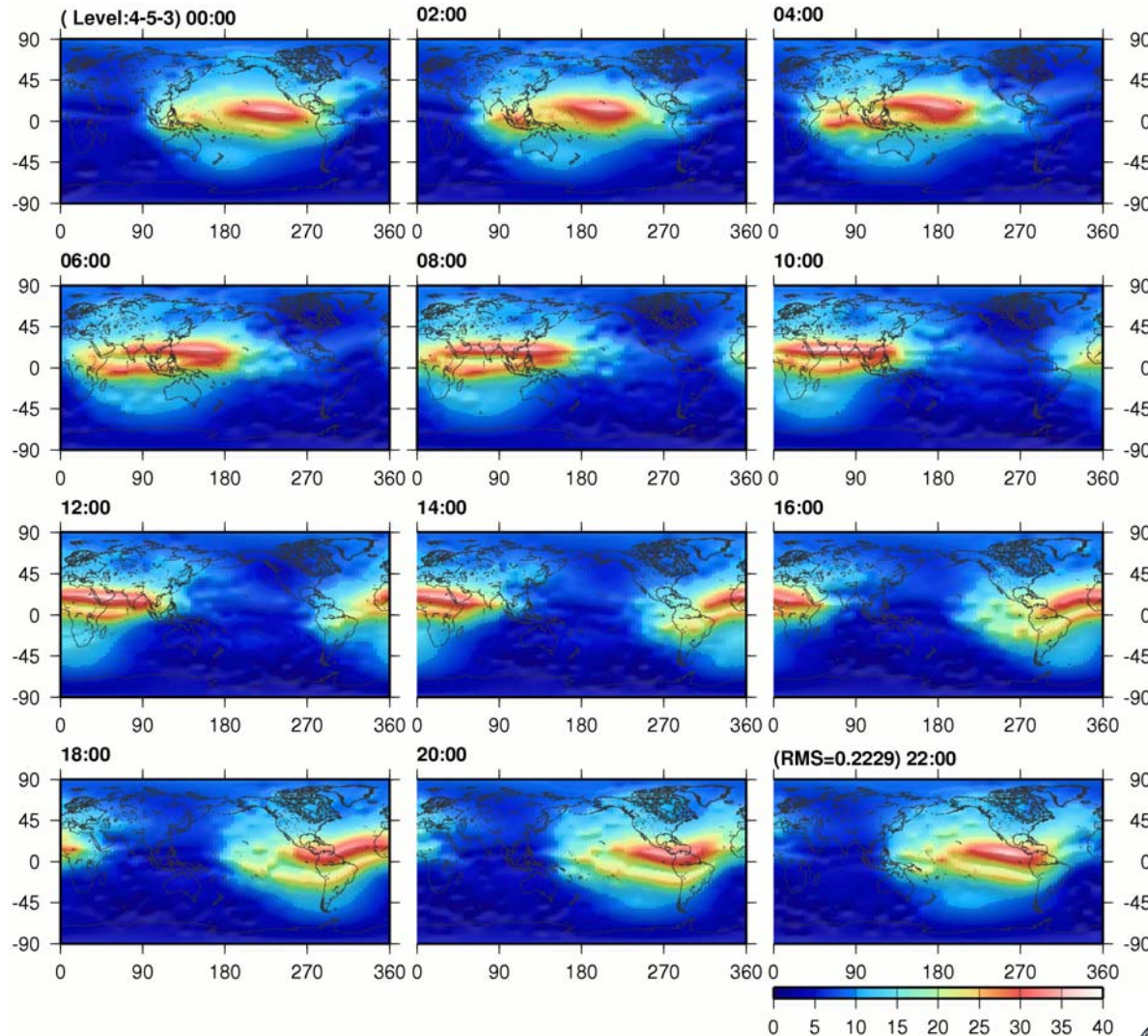
Estimated $\Delta VTEC$ 2006-07-21



- We apply this global approach for modeling $\Delta VTEC$, i.e. we introduce **trigonometric B-splines** in longitude (level 4).
- For modeling the latitude- and time-dependency we again use **endpoint interpolating B-splines** (level 5 and level 3, resp.).
- In the figures the crosses indicate the observation sites (blue and green colours mean negative values, orange positive values).
- Structures are modelable only in regions where observations are available.

COSMIC modeling — 3

Estimated VTEC 2006-07-21



- The panels show the **estimated VTEC** from COSMIC by adding the estimation of Δ VTEC to the reference values from IRI.
- The shown results are “**snapshots**” with a temporal resolution of 2 hours.
- Estimated VTEC values can be calculated **for each time t** of July 21st, 2006.
- **Validation**, e.g. by ionosondes necessary.

Final remarks

- We demonstrated how a **multi-dimensional** model of **ionospheric signals** can be performed.
- We presented first a **3D B-spline approach** with **time-dependent coefficients** for GNSS geometry-free observations.
- The accuracy was improved by incorporating **altimetry observations**.
- We further presented a **3D trigonometric/B-spline approach** for modeling VTEC data from COSMIC.
- We showed just some **components** of the procedure. The **next step** will focus on the **joint evaluation** of the real data from GNSS, altimetry and COSMIC.
- Since the presented procedure can be applied to **regional** or **local** areas, features like the **Equatorial Anomaly** can be handled appropriately.
- Finally, an **update of climatological parameters** of IRI can be performed.

